

and the Padda Island, than in the northern region, i.e. the Ongul Islands and the Langhovde District.

In conclusion, the ice sheet formerly extended all over these areas surrounding Syowa Station, and then its shrinking period began earlier in the northern part and later in the southern part, when cirque and valley glaciers were formed at one time on the higher part of the Langhovde District and a relative upheaval of land of about 15 to 20 meters took place. At present, these bared rock areas are a kind of "Antarctic oasis," and are being exposed to peri-glacial agencies and wind erosion. More detailed descriptions have been presented in following literatures.

Literature

- 1) Yoshikawa, T. and Toya, H. (1957): Preliminary report on geomorphological results of the Japanese 1956-57 Antarctic Research Expedition. Proceedings of IGU Regional Conference in Japan 1957, 240-242.
- 2) — (1957): Report on geomorphological results of the Japanese Antarctic Research Expedition, 1956-57. Antarctic Record, No. 1, 1-13 (in Japanese with English abstract).
- 3) — (1959): Glacio-geomorphological features of the neighbouring district of the Japanese Antarctic station, Syowa. (presented to the Antarctic Symposium of Buenos Aires in November 1959)

MOVEMENTS OF CALVED ICEBERGS ALONG PRINCE OLAV COAST*

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プリンスオラフ海岸における海岸線の変化*

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Concerning the icefront area along Prince Olav Coast, some amount of movements about a glacier tongue and calving icebergs were found out during two years.

For examining how much each of the calved icebergs floating on the sea or of the icebergs that are calving from glacier tongue has moved, photogrammetrical plotting was carried out to be compiled in a map of scale 1:50,000. The results are follows.

1. The mean distance of movement of calved icebergs is within 1200 to 800 meters in two years.

2. It is supposed that the drifting icebergs are not too much affected by either long-shore current or wind though limited in a narrow area.

3. Concerning the relative velocity of the stream flow of the icebergs, the central part of the glacier tongue is faster, while, the both sides are slower. The difference of the velocity between them might be caused by the ground feature beneath the glacier, and the larger the mass, especially about thickness, the faster it might moves.

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