## GEOMORPHOLOGICAL MAP OF EAST ONGUL AND WEST ONGUL ISLANDS

東オングル島および西オングル島地形学図 (1:10,000)

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Data sources: Fujiwara, K. (1971), Hayashi, M. (1977), Moriwaki, K. (1978) and Sawagaki, T. (1995)

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凡 例 LEGEND

2) 崖 Steps and Cliffs

海食崖 Cliffs by marine processes

地質および岩石構造制約による崖

流水による崖 Cliffs by fluvial, nival-fluvial processes

地形形態的要素 Morphography and Morphogenesis

Cliffs controlled by geology and lithologic structure

1) 傾斜変換線 Axes of curved slope segments, break of slopes

**EXPLANATION** 

et al., 1997; Hirakawa and Sawagaki, 1998; Miura et al., is released and deposited during the thawing season. 1998a, b, c). Geologically the Lützow-Holm Bay region is part of the Lützow-Holm Complex (Shiraishi et al., 1987) Submarine topography consisting of a diversity of rock types including: calcareous, N-S trending troughs and E-W trending basin-and-sill pelitic, quartzo-feldspathic, intermediate, basic and ultrabasic topography are remarkable in the submarine topography rocks. The Ongul Islands are part of the granulite-facies around East Ongul and West Ongul Islands. These terrain, and pyroxene gneiss is a common lithology. Garnet directions seem to be concordant with geological structures gneiss is dominant on East Ongul Island, while garnet-biotite such as the strike of foliation and the joint system of gneissic gneiss and granitic gneiss are dominant on West Ongul basement rocks as inferred from the geology of adjacent ice-Island. The geological (Ishikawa et al., 1994) and topo-free land areas. Stoss-and-lee topography is also present on graphic maps including submarine topography (Moriwaki, the sea floor. These submarine features are apparently 1979; Sawagaki, 1995a) are represented as separate small- drowned glacial landforms. There is a break of slope

Construction of the geomorphological map logical map: 1) Morphography and morphogenesis, 2) Geo- area is narrow or absent (Moriwaki, 1979). morphological processes, and 3) Individual geomorphological processes and their traces. Mapping of each feature is as accurate as possible according to the results of the field observations and aerial photographic interpretation. 1) Slopes and slope segments are depicted by

concave/convex breaks of slopes and by the axes of Antarctica. Nankyoku Shiryô (Antarct. Rec.), 41, 81-103. concave/convex slopes with gray solid/broken lines. Cliffs Fujiwara, K. (1973): The landforms of the Mizukumi Zawa near and steps are indicated by the tooth symbol with different Syowa Station, East Antarctica. Nankyoku Shiryô (Antarct. colors. Ridges and valleys, which are mostly controlled by Rec.), 46, 44-66. geological structure, are classified into three and two types Hayashi, M. (1977): The survey of submarine topography. using different colors and symbols. Although some ridges Report of JARE-16, 81-82. and small linear valleys are of glacial origin, they are all Hayashi, M. and Yoshida, Y. (1994): Holocene raised beaches represented as "structural" features by brown because of the difficulty of distinguishing them from structural ones. of slope, is not given.

areas with different colors: purple area for glacial process, Coast, East Antarctica -A report on a geomorphological green for marine process, pale-blue for fluvial and/or nival-survey during JARE-35 (1993-94)-. Nankyoku Shiryô fluvial process and yellow for eolian process. Glacial (Antarct. Rec.), 42, 151-167. landforms are classified as either depositional or erosional, Hirakawa, K., Ono, Y., Hayashi, M., Aniya, M., Iwata, S., and are depicted by light and dark purple, respectively. 3) Individual geomorphological processes and their traces are illustrated by purple symbols for erratics and glacial striae, ruby symbols for some periglacial landforms, blue

Igarashi, A., Harada, N. and Moriwaki, K. (1995): Marine symbols for nival-fluvial slope wash and red symbol for active honeycomb weathering. Fossil shells and abandoned penguin rookeries are represented by green symbols. Ishikawa, M., Shiraishi, K., Motoyoshi, Y., Tsuchiya, N.,

## Characteristics of main landforms

Glacial landforms The evidence of erratics and some glacial landforms indicates that the Ongul Islands were covered by the ice sheet. Glacial erosion by ice flow from the east scooped out

Miura, H., Maemoku, H., Igarashi, A. and Moriwaki, K. shallow depressions or troughs on bedrock trending ENE-WSW. Shallow depressions and rises trending N-S were also sculptured by the ice sheet, although they show

Bay. Spec. Map Ser. Natl Inst. Polar Res., 6 (with considerable rock control caused by dipping gneissic explanatory text 46p.). structure and different resistivities of various gneissic rocks. Miura, H., Maemoku, H., Seto, K. and Moriwaki, K. (1998b): Submarine depressions around the Ongul Islands also are Late Quaternary East Antarctic melting event in the Sôya glacial troughs eroded by N-S trending ice flow. The deep Coast region based on stratigraphy and oxygen isotopic ratio Ongul Strait, between the Ongul Islands and the continent, is of fossil molluscs. Polar Geosci., 11, 260-274. a N-S trending drowned glacial trough (Fujiwara, 1971). Miura, H., Moriwaki, K., Maemoku, H. and Hirakawa, K. Glacial striae remaining in a few places indicate two different (1998c): Fluctuations of the East Antarctic ice-sheet margin trends of N35°-65°E and N15°-45°E. Relatively thick till is distributed in some depressions.

## Marine landforms

Marine processes are recognized only in the area where Moriwaki, K. (1979): Submarine topography of the central part marine sediments are locally distributed. Raised beach of Lützow-Holm Bay and around Ongul Islands, Antarctica. deposits and marine mollusc fragments, however, can be Mem. Natl Inst. Polar Res., Spec. Issue, 14, 194-209. found up to 20 m a.s.l. in the Ongul Islands. TAMS- Omoto, K. (1977): Geomorphic development of the Sôya Coast, deposits are clearly classified into two groups of ca. 3.5-4.8 ka and ca. 30-46 ka. The older fossils are found in the Rep. Tohoku Univ., 7th Ser., 27, 95-148. Ongul Islands and the northern part of Langhovde, which are separated from the present ice sheet margin by a drowned heave activity in the vicinity of Syowa Station, East separated from the present ice sheet margin by a drowned glacial trough deeper than 500 m. On the basis of these Sawagaki, T. (1995b): Preliminary report on submarine facts, in addition to the deep continental shelf and a small topography around the Ongul Islands, East Antarctica. Proc. amount of isostatic uplift of Holocene raised beaches, NIPR Symp. Antarct. Geosci., 8, 213-218. Igarashi et al. (1995) and Miura et al. (1998a, b, c) reached Shiraishi, K., Hiroi, Y., Motoyoshi, Y. and Yanai, K. (1987): the following conclusions: 1) Marine transgression took Plate tectonic development of late Proterozoic paired place during the last interstadial period (Oxygen Isotope metamorphic complexes in eastern Queen Maud Land, East Stage 3 ?) and the Holocene. 2) Major deglaciation took place by the last interstadial period. 3) Expansion of the ice sheet during the LGM was slight although its extent is still unknown. 4) Sea-level during the last interstadial prior to the LGM was probably higher than that estimated from foraminiferal  $\delta^{18}$ O records in deep-sea sediments.

Periglacial landforms and phenomena

Tundra polygons, permafrost active-layer collapse and occurrence of periglacial creep/gelifluction are restricted in unconsolidated materials with sufficient soil water, such as

marine sediments and glacial tills. Measurement of ground The Ongul Islands lie approximately 5 km west of the temperature and frost heave near Lake Midori during 1992eastern coastline of Lützow-Holm Bay in Queen Maud Land, 1994 (Sawagaki, 1995b) revealed: 1) maximum thickness of East Antarctica. Among the many small islands, West the active layer of permafrost was 60-80 cm, 2) isotherm Ongul and East Ongul Islands are the largest, with low-lying lines of 0°C suggested that seasonal freezing during late undulating features below 48 m a.s.1. Syowa Station, the summer had occurred both downward from the ground main scientific station of Japanese Antarctic Research surface and upward from the permafrost table, 3) frequent Expeditions, is situated in the northern part of East Ongul diurnal frost heaves were caused by freeze-thaw cycles above Island. Since 1957 a number of studies have been done on 10 cm depth, and 4) seasonal frost heave was 3 mm. Frost the raised beach deposits and glacial landforms along the ice-shattering occurs frequently in restricted locations where free rocks of the Sôya Coast facing Lützow-Holm Bay meltwater from drift snow runs during the austral summer. (Yoshikawa and Toya, 1957; Fujiwara, 1973; Moriwaki, Nival-fluvial are surface wash and rill erosion processes 1976; Omoto, 1977; Yoshida, 1983; Hirakawa et al., 1984; operating around the drift snow. The niveo-eolian process Hayashi and Yoshida, 1994; Igarashi et al., 1995; Maemoku takes place in localities where eolian sand on the drift snow 68°59′20″S

between the submarine and subaerial topographies on East Ongul and West Ongul Islands. The slope above sea level is gentle while the slope under water is steep. Such a break There are three main legend groups on the geomorpho- is not as conspicuous on adjacent small islands where a flat

## Fujiwara, K. (1971): Soundings and submarine topography of

the glaciated continental shelf in Lützow-Holm Bay,

in the Lützow-Holm Bay region, East Antarctica. Holocene Environmental Changes in Antarctic Coastal Areas, ed. by P.A. Berkman and Y. Yoshida, Mem. Natl Inst. Polar Res., Spec. Issue, 50, 49-84. Hirakawa, K. and Sawagaki, T. (1998): Radiocarbon dates of 2) Specific geomorphological processes are indicated by fossil shells from raised beach sediments along the Sôya

> Fujiwara, K., Moriwaki, K. and Yoshida, Y. (1984): Antarctic geomorphological map of Langhovde. Spec. Map Ser. Natl Inst. Polar Res., 1 (with explanatory text 63p.). fossils of 30-40 ka in raised beach deposits, and late Pleistocene glacial history around Lützow-Holm Bay, East

Selected radiocarbon dated ages of fossil shells are given on Shimura, T. and Yanai, K. (1994): Geological map of Ongul Islands. Antarct. Geol. Map Ser., Sheet 36 (with explanatory Maemoku, H., Miura, H., Saigusa, S. and Moriwaki, K. (1997): Stratigraphy of the late Quaternary raised beach deposits in the northern part of Langhovde, Lützow-Holm Bay, East Antarctica. Proc. NIPR Symp. Antarct. Geosci., 10, 178-

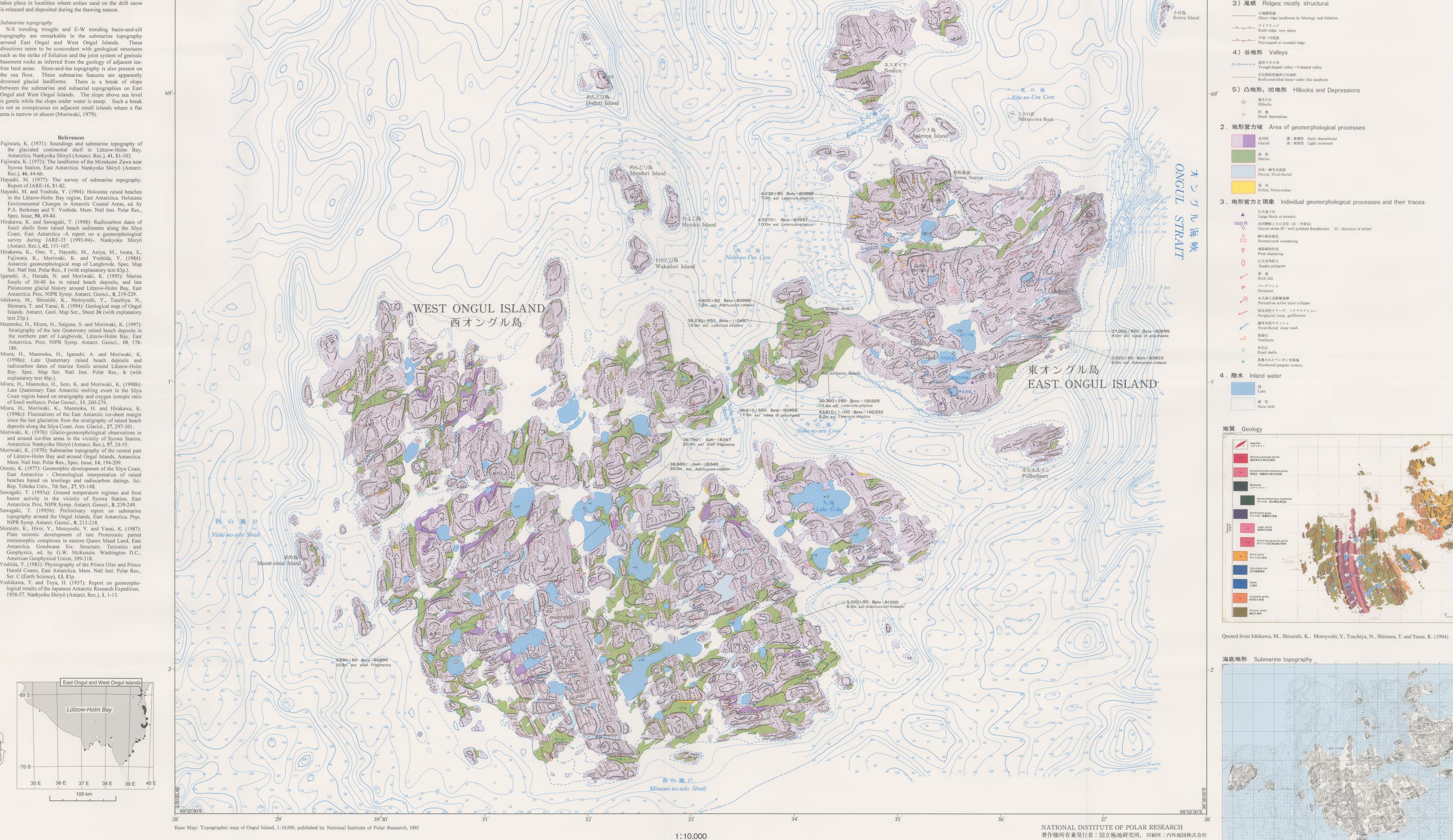
> (1998a): Late Quaternary raised beach deposits and radiocarbon dates of marine fossils around Lützow-Holm

since the last glaciation from the stratigraphy of raised beach deposits along the Sôya Coast. Ann. Glaciol., 27, 297-301. Moriwaki, K. (1976): Glacio-geomorphological observations in and around ice-free areas in the vicinity of Syowa Station, Antarctica. Nankyoku Shiryô (Antarct. Rec.), 57, 24-55. beaches based on levelings and radiocarbon datings. Sci.

Antarctica. Gondwana Six: Structure, Tectonics and Geophysics, ed. by G.W. McKenzie. Washington D.C., American Geophysical Union, 309-318.

Ser. C (Earth Science), 13, 83p. Yoshikawa, T. and Toya, H. (1957): Report on geomorphological results of the Japanese Antarctic Research Expedition, 1956-57. Nankyoku Shiryô (Antarct. Rec.), 1, 1-13.

Lützow-Holm Bay



メルカトール図法(Lat69°01')