

# Application of isotope analysis of nitrate and sulfate for ice cores and polar environments

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Isotope analysis is generally known useful tool to identify sources and formation pathways for atmospheric trace species. So far, isotope analysis of  $\text{SO}_4^{2-}$  and  $\text{NO}_3^-$  preserved in ice core at Summit, Greenland, anthropogenic impact for atmospheric sulfur dioxide ( $\text{SO}_2$ ) and nitrogen oxides ( $\text{NO}_x = \text{NO} + \text{NO}_2$ ) have been reported (e.g. Patris et al., 2000; Hasting et al., 2009). In addition to conventional isotopes of  $\delta^{15}\text{N}$  and  $\delta^{34}\text{S}$ , triple oxygen isotope compositions ( $\Delta^{17}\text{O}$ ) and quadruple sulfur isotope compositions ( $\Delta^{33}\text{S}$  and  $\Delta^{36}\text{S}$ ) are new atmospheric finger prints which enable us to reconstruct how atmospheric reactions work/worked in the past and present atmosphere. In this presentation, we introduce general concepts, methods, and recent our studies applying isotope analysis of  $\text{SO}_4^{2-}$  and  $\text{NO}_3^-$  to polar environments.

## 1. Theory and concepts

With two, three and four isotopes for N, O, S respectively, a total six isotope ratios can be considered ( $\delta^{15}\text{N}$ ,  $\delta^{17}\text{O}$ ,  $\delta^{18}\text{O}$ ,  $\delta^{33}\text{S}$ ,  $\delta^{34}\text{S}$ ,  $\delta^{36}\text{S}$ ). Thus,  $\text{SO}_4^{2-}$  and  $\text{NO}_3^-$  are isotopically defined by three and four  $\delta$  values, respectively. However, most of chemical and physical processes follow mass-dependent fractionation (MDF), following relationship hold for the oxygen and sulfur isotope systems.

$$\delta^{17}\text{O} = 0.52 \times \delta^{18}\text{O}$$

$$\delta^{33}\text{S} = 0.515 \times \delta^{34}\text{S}$$

$$\delta^{36}\text{S} = 1.91 \times \delta^{34}\text{S}$$

Thus, it had been unnecessary to measure all six  $\delta$  values, as they were thought not to be providing useful information for long time. However, since Mark Thiemens and his colleagues at University of California in San Diego have found mass-independent fractionation (MIF) for oxygen in  $\text{O}_3$  formation (Thiemens and Heidenreich, 1983) and sulfur in Archean Rock (Farquhar et al., 2000), and these new definitions of deviation from MDF are defined as following formulas and started to apply to various geochemical studies.

$$\Delta^{17}\text{O} = \delta^{17}\text{O} - 0.52 \times \delta^{18}\text{O}$$

$$\Delta^{33}\text{S} = \delta^{33}\text{S} - 0.515 \times \delta^{34}\text{S}$$

$$\Delta^{36}\text{S} = \delta^{36}\text{S} - 1.91 \times \delta^{34}\text{S}$$

Atmospheric ozone ( $\text{O}_3$ ) possesses high  $\Delta^{17}\text{O}$  values of approximately 26‰ in contrast to most of the oxygen bearing compounds such as  $\text{O}_2$  and  $\text{H}_2\text{O}$  which have  $\Delta^{17}\text{O}$  values of approximately 0‰. The oxygen atoms of  $\text{O}_3$  are directly or indirectly transferred to  $\text{SO}_4^{2-}$  and  $\text{NO}_3^-$  through the various oxidation pathways of their precursors,  $\text{SO}_2$  and  $\text{NO}_x$ . For sulfur isotopes, a series of groundbreaking studies has shown that sulfur isotopes in sulfate from Plinian eruptions show MIF (e.g. Baroni et al., 2007), but background stratospheric sulfate aerosols from atmospheric carbonyl sulfide (OCS) and its breaking down OCS mainly photolysis, show no evidence of sulfur MIF (Hattori et al., 2011; Schmidt et al., 2013). Consequently,  $\Delta^{17}\text{O}$ ,  $\Delta^{33}\text{S}$  and  $\Delta^{36}\text{S}$  are new proxy for the ice core and polar environmental studies.

## 2. Measurement methods

### 2.1 Nitrate

Previously, the measurements of nitrogen isotope and oxygen isotopes are needed separately, new approach that uses the generation of  $\text{N}_2\text{O}$  from  $\text{NO}_3^-$  reduction enable us to measure  $\delta^{15}\text{N}$ ,  $\delta^{18}\text{O}$ , and  $\Delta^{17}\text{O}$  simultaneously. Recently we developed automated method measuring  $\delta^{15}\text{N}$ ,  $\delta^{18}\text{O}$ , and  $\Delta^{17}\text{O}$  simultaneously by using the bacterial method (Sigman et al., 2001) coupled with  $\text{N}_2\text{O}$  decomposition via microwave-induced plasma (MIP) (Hattori et al., 2016).

### 2.2 Sulfate

In contrast to  $\text{NO}_3^-$ , isotope analysis of oxygen and sulfur for  $\text{SO}_4^{2-}$  need to be performed separately. For  $\Delta^{17}\text{O}$  measurement of  $\text{SO}_4^{2-}$ , first  $\text{SO}_4^{2-}$  is separated from other ions by ion chromatographic method (Geng et al., 2013) or anion-exchange resin method (Le Gendre et al. 2017). Separated  $\text{SO}_4^{2-}$  chemically converted to silver sulfate, and this  $\text{Ag}_2\text{SO}_4$  powder is converted to  $\text{O}_2$  with high temperature (1000°C) for isotope measurements (see Geng et al., 2013; Ishino et al., 2017 for more detailed protocols). For  $\Delta^{33}\text{S}$  and  $\Delta^{36}\text{S}$ ,  $\text{SO}_4^{2-}$  is converted to  $\text{SF}_6$  by fluorination for measuring  $\text{SF}_5^+$  ions of m/z 127, 128, 129, 131.

## 3. Application for Antarctic study

Some studies have applied above methods for Antarctic ice cores (e.g. Sofen et al., 2014), but information of temporal and spatial variations of atmospheric  $\text{SO}_4^{2-}$  and  $\text{NO}_3^-$  are limited. In particular, there remain many theoretical assumptions for the controlling factors of  $\Delta^{17}\text{O}$  of  $\text{SO}_4^{2-}$  and  $\text{NO}_3^-$  in those model estimations. To test one of those assumption that  $\Delta^{17}\text{O}$  of  $\text{O}_3$  have a flat value and do not influence the seasonality of  $\Delta^{17}\text{O}$  values of  $\text{SO}_4^{2-}$  and  $\text{NO}_3^-$ , we performed the first simultaneous measurement of  $\Delta^{17}\text{O}$  values of  $\text{O}_3$ ,  $\text{SO}_4^{2-}$  and  $\text{NO}_3^-$  collected at Dumont d'Urville (DDU) Station, Antarctica. As a result,  $\Delta^{17}\text{O}$  values of  $\text{SO}_4^{2-}$  and  $\text{NO}_3^-$  exhibited seasonal variation characterized by minima in the austral summer and maxima in winter, while  $\Delta^{17}\text{O}$  values of  $\text{O}_3$  showed no significant seasonal variation. These contrasting seasonal trends suggest that seasonality in  $\Delta^{17}\text{O}$  values of atmospheric  $\text{SO}_4^{2-}$  and  $\text{NO}_3^-$  is not the result of changes in  $\Delta^{17}\text{O}$  values of  $\text{O}_3$ , but of the changes in oxidation chemistry (Ishino et al., 2017). We have also measured  $\delta^{34}\text{S}$ ,  $\Delta^{33}\text{S}$  and  $\Delta^{36}\text{S}$  for the same samples, and found seasonal pattern of  $\delta^{34}\text{S}$ , showing not only biogenic sulfur source but also volcano and stratospheric sulfur sources might contribute in winter. In

addition, no significant deviation of sulfur MIF is found, suggesting different mechanisms worked for atmospheric  $\text{SO}_4^{2-}$  production of polluted air in which small sulfur MIF was typically observed.

Conversely, we also found the annual mean  $\Delta^{17}\text{O}$  values for atmospheric  $\text{SO}_4^{2-}$  at coastal Antarctica (Ishino et al., 2017) is not matched with the  $\Delta^{17}\text{O}$  values preserved in the inland Antarctic ice core records (e.g. Alexander et al., 2002; Sofen et al., 2014). In order to test this, spatial variation of isotopic compositions between coastal site and inland in eastern Dronning Maud Land, East Antarctica collected during the 54th and 57<sup>th</sup> Japanese Antarctic Research were performed.  $\Delta^{17}\text{O}$  values of nss- $\text{SO}_4^{2-}$  at the East Antarctica ranges from 2.2 to 3.3‰, and the  $\Delta^{17}\text{O}$  value of nss- $\text{SO}_4^{2-}$  for coastal site was lower than those for inland site. Thus, this result suggested that oxidizing chemistry for biogenic sulfur is different among coastal and inland sites, although small sulfur isotopic variations are observed and source of sulfur is biogenic and homogeneous (Uemura et al., 2016). For the isotopic compositions of  $\text{NO}_3^-$ , considerably increasing values of  $\delta^{15}\text{N}$  of  $\text{NO}_3^-$  are observed from coastal to inland sites. The  $\delta^{18}\text{O}$  and  $\Delta^{17}\text{O}$  of  $\text{NO}_3^-$  values, on the other hand, decreases with increasing of  $\delta^{15}\text{N}$  values, indicating the secondary formation of  $\text{NO}_3^-$ . Thus, spatial variations of isotopic compositions of  $\text{NO}_3^-$  reflect the post-depositional processes on the East Antarctic snow, as reported widely in Antarctica (Erbrand et al., 2013; Shi et al., 2016).

#### 4. Application for Arctic study

In previous study, the decrease in  $\delta^{15}\text{N}$  of  $\text{NO}_3^-$  is strongly correlated with fossil fuel emissions estimates since 1750 (Hasting et al., 2009), showing  $\delta^{15}\text{N}$  as potential tool for reconstruction of past natural and anthropogenic  $\text{NO}_x$  emission. For  $\Delta^{17}\text{O}$  in ice-core  $\text{NO}_3^-$  is also reported to provide information on past ratio of  $\text{O}_3/(\text{HO}_2 + \text{RO}_2)$  (Geng et al., 2017). Certainly, nitrogen and oxygen isotopic compositions of  $\text{NO}_3^-$  provide information on changes in the nitrogen source and its formation pathways, but ice core records for  $\text{NO}_3^-$  concentrations and its isotopic compositions are problematic because of post depositional loss of  $\text{NO}_3^-$  via photolysis. Although Greenland generally have higher accumulation of snow than that in Antarctica, but post-depositional loss of  $\text{NO}_3^-$  in ice cores in Greenland is still controversial. We analyzed isotopic compositions of  $\text{NO}_3^-$  preserved in the high-accumulation dome ice core, South East Greenland (SE-Dome), which has a dome with high accumulation rate (about  $1 \text{ m yr}^{-1}$ ) drilled by Iizuka and his colleagues (Iizuka et al., 2016). The nitrogen isotopic compositions for  $\text{NO}_3^-$  were generally lower than those reported in Summit, Greenland, suggesting that some extent of  $\text{NO}_3^-$  deposited in Summit was removed via photolysis. We found decreases of  $\delta^{15}\text{N}$  of  $\text{NO}_3^-$  in SE-Dome, but no significant changes for  $\Delta^{17}\text{O}$  during past 60 years.

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